Dynamical Climatology

Aspects of dynamics of the middle atmosphere inferred using data from a satellite and from a numerical model.

by

T.D.A. Fairlie and A. O'Neill

DCTN 54

April 1987

Meteorological Office (Met. O. 20)
London Road
Bracknell
Berkshire RG12 2SZ

DYNAMICAL CLIMATOLOGY TECHNICAL NOTE NO. 54

ASPECTS OF DYNAMICS OF THE MIDDLE ATMOSPHERE INFERRED USING DATA FROM A SATELLITE AND FROM A NUMERICAL MODEL

by

T.D.A.Fairlie and A.O'Neill

This paper has been submitted to Phil. Trans. Roy. Soc. for inclusion in the proceedings of a meeting on Studies of the Middle Atmosphere, held at the Royal Society 4-5 December 1986.

Met O 20 (Dynamical Climatology Branch)
Meteorological Office
London Road
Bracknell
Berkshire RG12 2SZ U.K.

April 1987

Note: This paper has not been published. Permission to quote from it should be obtained from the Assistant Director of the above Meteorological Office Branch.

ABSTRACT

The data used in this study are from a Stratospheric Sounding Unit (SSU) on a polar-orbiting satellite, and from a numerical model of the middle atmosphere based on the primitive equations. They are used to discuss the dynamics of the major stratospheric warming of 1984/85. The warming is of note because the westerly circulation started to break down rapidly when the westerly vortex was larger than might be expected of a state "pre-conditioned" for a major warming. At the height of the warming, there is a sudden loss of high values of Ertel's potential vorticity on isentropic maps derived from the SSU. Ground based measurements and a simulation of the warming with the numerical model indicate that strong gradients of temperature arise in westward-sloping "fronts," which the SSU is unable to resolve fully. The simulated stratospheric fronts have features in common with observed fronts that form in the upper troposphere, though there may turn out to be dynamical differences between the two cases.

1. INTRODUCTION

Data from satellites are greatly improving our knowledge of the middle atmosphere. They have revealed a wealth of flows evolving on a wide range of scales in space and time. Before such data came along, information on the region came mainly from radiosondes and rocketsondes. A long and valuable sequence of hand-drawn analyses of the lower stratosphere of the northern hemisphere has been made daily from these sources by workers at the Free University of Berlin. But the coverage of data is not uniform, being extremely sparse over the oceans and in the southern hemisphere. Moreover, measurements are not made with instruments of equal reliability. Instruments on satellites, on the other hand, can make measurements globally and with consistent quality. These attributes are especially important for diagnostic studies of dynamics (measured fields often have to be differentiated), and for initialising and validating simulations with numerical models.

As more complex flows have been observed, our dynamical thinking about the middle atmosphere has changed. Many diagnostic studies have been based on the theory of wave, mean-flow interaction. In a typical application (e.g., Palmer, 1981), a wave is taken as the departure from the current zonal-mean state, and is supposed to propagate quasi-linearly on that state. It is now becoming recognised that when the middle atmosphere is strongly perturbed a partition into a wave and a mean flow may not be useful. A return to a synoptic description has been prompted by recent papers by McIntyre and Palmer (1983, 1984, 1985). They emphasised the dynamical importance of Ertel's potential vorticity, Q, and based their studies of the stratosphere on isentropic maps of Q derived from measurements by a stratospheric sounding unit (SSU) on board the satellite, TIROS-N, of NOAA's operational series.

A good approximation to Q for large-scale circulations is given by

$$Q \approx -g(\zeta + f)\,\partial\theta/\partial p \qquad (1),$$

where g is the acceleration due to gravity, ζ is the relative vorticity, f is the Coriolis parameter, θ is potential temperature (constant θ defines an isentropic surface), and p is pressure. Q has two important properties for a dynamical analysis. (1) For adiabatic, inviscid flow, it is conserved along an air parcel's trajectory, as is θ . Contours of Q on isentropic maps are then material lines which can be used to follow the movement of air. In the mid stratosphere, this can be done for only a week or so at a time, owing to the non-conservative effects of radiation; in the upper stratosphere, the limit is a few days. (2) Under balanced conditions (e.g. thermal wind balance) and with boundary conditions, all other fields can, in principle, be derived from the three-dimensional distribution of Q. Hoskins et al. (1985) elaborate on these properties and give examples of their use. The uniform quality of satellite data is an advantage in calculating Q because measurements of temperature or thickness must be differentiated twice (first to obtain winds, then to obtain relative vorticity).

Satellite-borne instruments do not take readings at a point in the atmosphere but average over a volume of it. Thus they yield only a somewhat blurred view of the distribution of Q, and smaller-scale features embedded in the large-scale flow are smoothed or lost altogether. As we shall see, sharp changes in the variation of temperature with height, captured in radiosonde measurements, are smeared out; and narrow tongues of material which growing anticyclones draw from the westerly vortex are inadequately resolved, so the transport of material is obscured.

Failure to resolve such features in a numerical model could be detrimental to the model's ability to reproduce accurately even the large-scale flow of the middle atmosphere, especially in long integrations. Resolution in the middle atmosphere comparable with or, indeed, exceeding that now used by general circulation or forecasting models of the troposphere could well be needed.

In this paper, we give an example of how data from a satellite can be used in conjunction with a numerical experiment to learn about the dynamics of a tantalising phenomenon that has yet to be fully understood--the stratospheric sudden warming. The first major warming to be studied extensively with such data happened in the northern hemisphere in February 1979 (e.g. Palmer, 1981). We focus on a major warming which took place in mid winter, 1984/85. In this spectacular event, the westerly vortex was split into two roughly equal parts. Scales of motion arise that SSU measurements do not adequately resolve, as is signalled by a sudden loss of high values of Q at the centre of the newly-formed vortices. We use a numerical model of the middle atmosphere with a prescribed lower boundary to simulate the warming. The model has better spatial resolution than our observations, and we find that narrow, westward-tilting zones with strong temperature gradients develop in the stratosphere. We refer to them as "fronts." The atmospheric structure near these fronts is reminiscent of that found near fronts that have been observed in the upper troposphere.

The plan of our paper is this: In section 2, we describe the data and its method of analysis, after which we summarize the features of our numerical model. The observations and the model are used in section 3 to discuss aspects of the dynamics of the major warming of 1984/85. Concluding remarks are in section 4.

2. DESCRIPTION OF DATA AND NUMERICAL MODEL

(a) Data and method of analysis

We use data mainly from a stratospheric sounding unit (SSU) on board the satellite, NOAA-7, of the TIROS-N series. Pick and Brownscombe (1981) describe the instrument. The satellite is in sun-synchronous orbit at a height of 850km, making about 14 orbits a day. Stratospheric analyses are made from radiances measured by the SSU, by a microwave sounding unit (MSU), and by a high-resolution infrared sounder (HIRS-2). All three are nadir sounding radiometers which scan across the orbital path giving almost global coverage with good horizontal resolution. The measuring channels of the SSU have weighting functions centred in the stratosphere near 15, 5, and 1.5 mb. Statistical tests show that the vertical resolution of measurements is about 10km. Resolution is somewhat poorer for the SSU providing data for this study because the channel for 5mb was inoperative. The performance of the system was not seriously degraded, however. We compared analyses with those derived from an SSU on the satellite, NOAA-6, and found that differences could be neglected for our purposes. (Data from NOAA-6 were received only sporadically, and its HIRS instrument was inoperative at the time.)

Details of the calculation of isentropic maps of potential vorticity are given by Clough et al.(1985), who also discuss the errors involved. Briefly, thicknesses are derived from radiances by statistical regression. They are then added daily to a global analysis of geopotential height at 100mb made at the Meteorological Office as part of an operational forecasting cycle. Fields of geopotential height in the stratosphere are thus obtained at 20, 10, 5, 2 and 1

mb, and they are supplemented by an operational analysis at 50 mb. Potential vorticity is calculated by differentiating and interpolating these fields.

(b) The numerical model

We use a global model of the stratosphere and mesosphere based on the primitive equations. The height of its lowest isobaric surface, 100mb, is prescribed using observations. A full description of the model was given by Butchart et al. (1982), though our version uses finite differencing accurate to fourth order rather than to second order. Summarising its main features, it has a regular grid in spherical co-ordinates with grid points at intervals of 5° in latitude and longitude, and 33 levels equally spaced in log pressure, about 2.5km apart. In the experiment discussed here, radiative processes are approximated by Newtonian cooling, whereby temperature is reduced towards its local value in a reference state (the zonally averaged state at the start of an experiment) on a time-scale of about 20 days in the lower and middle stratosphere and of about 5 days in the upper stratosphere. Drag due to the dissipation of gravity waves is represented using Holton's (1976) parametrization of Rayleigh friction.

Initial data needed to start an integration with the model are obtained for the stratosphere from the analyses described above. For the mesosphere, we used a climatology prepared using satellite data at the University of Oxford (courtesy of J J Barnett and M Corney).

3. THE MAJOR WARMING OF 1984/85

(a) Observations

A dramatic and unusually early major warming developed at the end of December 1984 in the stratosphere of the northern hemisphere. A summary of its main synoptic features is given by Labitzke et al. (1985). Fig. 1(a) shows the zonally averaged, geostrophic wind on 20 December 1984. There is a broad jet with winds exceeding 80 ms^{-1} in the upper stratosphere. By 2 January 1985, Fig.1(b) shows that westerlies have been replaced by easterlies in the zonal mean throughout the stratosphere at high latitudes. Near 1 mb, zonal-mean easterlies have speeds greater than 70 ms^{-1}, and locally temperatures rise to about 20°C in the mid stratosphere according to measurements made using lidars at Haute Provence and at the University of Aberystwyth (Labitzke et al., Fig.7).

It has been argued that major warmings such as this one are favoured after the stratosphere has been "pre-conditioned" (e.g. Kanzawa, 1984). That is, the stratosphere attains some state, perhaps as a result of an earlier minor warming, that lends itself to disruption by a large-scale, tropospheric disturbance. Such a state has been taken to be one in which the zonal-mean jet in the stratosphere is narrow and confined to high latitudes. The argument is based on the theory of wave, mean-flow interaction. Since stationary waves propagate only where the winds are westerly (Charney and Drazin, 1961), the theory predicts that they will be focused by a narrow jet into high latitudes and that a strong polar warming and rapid deceleration of zonal winds are thereby favoured. McIntyre and Palmer (1983) extended this idea to include flows that are asymmetric about the pole. They predict that an unusually small

westerly vortex in the stratosphere will typify conditions prior to a major warming.

The major warming of 1984/85 is surprising in that the stratosphere was not pre-conditioned in either way. The flow depicted in Fig.1(a), at the start of rapid deceleration of the zonal-mean wind, comprised not only a broad zonal-mean jet but also a westerly vortex that was not unusually small for the time of year, as we have verified using SSU data for 8 years. Moreover, although in the troposphere the large-scale disturbances driving the warming were strong (as measured by variations of geopotential height around latitude circles), they were not much stronger than disturbances which develop in winters not having major warmings. During the build-up to a major warming the flow is too disturbed for the linear approximation to be valid. We conclude that predictions about warmings are unreliable when they are based on the supposed quasi-linear propagation of planetary waves. O'Neill and Pope (1987) discuss this point further.

The movement of material during the warming can be followed, albeit with a blurred view, using isentropic maps of potential vorticity, Q. Fig. 2 is such a map for 22 December 1984 and for the 850K isentropic surface which lies near 10mb, at a height of about 30Km in the mid stratosphere. Arrows representing the speed and direction of the winds on this surface are also shown. There is a strong anticyclonic circulation, associated with the so-called Aleutian High, centred on the date-line at mid latitudes. Air in the anticyclone has low values of Q, characteristic of air from sub-tropical latitudes. On the map, there are two patches of low Q which, according to the winds near them, would circulate around the anticyclone. The location of individual contours of Q are not known reliably near the anticyclone, however, because gradients of Q are weak, and a small error in calculating Q (caused, perhaps, by poorly resolved vertical

structure) can make a big difference to a contour's position. The main westerly vortex is highly elliptical, and there is an intense jet stream across the polar cap where the anticyclone abuts the cyclone. This figure illustrates our earlier remark: it would be hard to justify a dynamical description of the ensuing warming which took the zonal mean of such a state as a "basic state" for the propagation of planetary waves. For such a strongly disturbed flow, O'Neill and Pope (1987) advocate a description based on vortices interacting with each other rather than one based on waves interacting with a mean flow (though the theory of vortex interactions, developed for example by Dritschel, 1986, is still rudimentary for stratospheric applications).

Fig.3 shows the circulation and distribution of Q on 26 December 1984. The westerly vortex is now even more elongated and there are two closed cyclonic circulations within it centred on local maxima in Q. An anticyclone has developed over the Atlantic, and the associated circulation is drawing out a tongue of high Q into low latitudes. Such advection of material by developing anticyclones is common in the winter stratosphere. Its occurrence was first pointed out by McIntyre and Palmer (1983). They regard such irreversible (time oriented) advection of Q as being the hallmark of breaking planetary waves though, as O'Neill and Pope demonstrate, the phenomenon is not restricted to cases for which a propagating wave is readily defined. The area covered by the westerly vortex shrinks as high Q is drained away by growing anticyclones. We are still assuming in interpreting Fig.3 that Q contours are approximately material lines. This assumption gets poorer as the warming develops for two reasons: (1) Temperatures rise in the upper stratosphere and warm air overrides cold air; the vertical resolution of the SSU is then a limitation.
(2) Temperatures are far from their values at radiative equilibrium so air cools more quickly and Q is not conserved for so long.

Eventually, the westerly vortex splits, as shown in Fig.4 for 30 December 1984. Air from low latitudes streams over the pole between the two cyclones. The maximum values of Q ought to be conserved for a few days at least, though advection might move the maxima somewhere else on an isentropic map. Fig.5 for 2 January 1985 shows, however, that peak values are halved for one of the vortices and are almost so for the other. The reason is that there are now sharp changes in vertical temperature gradient which the SSU cannot resolve, and the vertical derivative in Eq.(1) is accordingly underestimated. Fig.6 shows the variation of temperature with height (solid curve) recorded by a radiosonde launched from Berlin on 2 January. Whereas the lower stratosphere is nearly isothermal, there is a highly stable layer in the mid stratosphere where temperatures increase with height at about 10K/km. The dashed curve on the figure is computed from analyses of geopotential height derived using SSU data. The profile is colocated with the Berlin radiosonde. The sharp transition to a stable layer is not captured; the stability of the mid stratosphere is underestimated, and so is Q. This figure clearly exposes the limitations of data from satellites on such occasions-- but there is mitigation. We find with Clough et al. (1985) that Q derived from SSU data is usually well conserved from day to day. So when it is not, we are warned that our analysis is poorly representing the state of the stratosphere.

With the split of the vortex, an anticyclonic circulation enclosing air drawn from low latitudes (as indicated by low Q) encompasses the polar cap (Fig.5). The two cyclones are being stretched around it. Because the easterly winds around the anticyclone increase in speed with height, we might expect that the cyclones would move westward faster at 1mb than at 10mb, i.e. that the cyclones would lean over. This is confirmed by our analyses, with the caution that they are not as reliable at this stage of the warming as they are at its

inception. We ascribe the loss of resolution in the vertical to a generation of structure by a combination of such differential advection and adiabatic warming as air sinks. (Without vertical motion, temperatures could not attain the values observed.)

Satellite measurements do not reveal fully the strong temperature gradients and wind shears that develop during major stratospheric warmings. A numerical simulation that we made of the warming of 1984/85 reproduces strong vertical temperature gradients that are more in keeping with the observations made by lidars and radiosondes. We are able to discern what we call "fronts" in the stratosphere. Their three-dimensional structure would not be well exposed even by lidars and radiosondes because there are few such measurements in the middle and upper stratosphere. We now turn to a description of the simulated circulation.

(b) Simulation with a numerical model

The simulation of the major warming was initialised using observational data for 22 December 1984, and the lower boundary of the model was updated daily with analyses of geopotential height at 100mb. In the simulated stratosphere, the split in the westerly vortex was successfully reproduced, the elongating vortex having about the correct orientation but splitting a day late.

Fig.7 shows the simulated distribution of Q and winds on the 850K surface on 30 December 1984, just before the vortex splits. Contours of Q, or material lines, are being irreversibly deformed. The bottom half of the figure shows a long tongue of high Q drawn around an anticyclone, which itself is centred on a patch of low Q injected from low latitudes. The top half of the figure shows another long tongue, this time of low Q, in the process of wrapping around the other anticyclone.

Fig.7 shows clearly why it may prove necessary to use models with high resolution in order to simulate well the seasonal evolution of the middle atmosphere. Tongues of low or high Q get narrower as they lengthen and fail to be resolved by our model. They may vanish through numerical truncation. Such truncation (and other numerical devices used to keep an integration stable) affects small scales of motion more than large scales. In flows containing highly deformed Q contours, the simulated circulation might therefore be damped more than it should be because Q is not conserved well enough. Such flows are common in the stratosphere during winter, especially in the northern hemisphere. Although our model (5°x5° grid boxes) simulates well the development of a stratospheric warming, the flow observed after a warming is not reproduced so well, even when we use a radiation scheme that is thought to be reasonably accurate (provided by K. Shine). Part of the explanation is probably that small

scales of motion which develop in the horizontal and in the vertical are inadequately resolved.

Further examples of small-scale features are the "fronts" that form in the stratosphere at the height of the simulated warming. Fig.8 is a cross-section of temperature near 60°N on 31 December 1984. (Figs.7 and 8 are chosen one day apart to highlight particular features.) There are two zones where temperatures change rapidly with height and longitude. These strongly baroclinic zones, or fronts, slope westward with height and are embedded in a flow having a variation in temperature on a much larger scale--mainly in wavenumber 2 around a circle of latitude. Below a front, the stratosphere is approximately isothermal, whereas in a front temperatures increase with height at about 10K/km, variations in good agreement with those found from observations (Fig.6, solid curve). Between the fronts lie deep isothermal layers where the stratopause (the level where temperatures start to decrease with height) is not well marked.

The steep gradients in temperature are associated with strong shears in winds. There are strong poleward jets to the east of the fronts and weaker equatorward flow just to their west (this may be inferred for the mid stratosphere by locating the temperature fronts on Fig.7). The fronts are also well marked by steep gradients in relative vorticity (not shown). The magnitude of relative vorticity about equals that of planetary vorticity (the Coriolis parameter) near the front. Ageostrophic motions must therefore be important, as is confirmed by the large rises in temperature. They far exceed what could be achieved by warm advection by the geostrophic winds. Air warms as much as it does because it sinks near the front (at up to 15 cm s^{-1} in the plane represented by Fig.8), and isentropic surfaces dip there as a result.

A cross-section of ln(Q) is shown in Fig.9. (Taking the logarithm gives an even distribution of contours.) There are folds in the contours near the fronts. Near 120°W, a downward-pointing tongue of high Q lies next to an upward-pointing tongue of low Q. The fold does *not* develop solely by a two-dimensional circulation in height and longitude, though air is indeed sinking near the protrusion of high Q and rising near that of low Q. Contours of Q on the west side of the fold coincide with isentropic surfaces (not shown) and dip as air sinks. The tongue of low Q does not coincide with isentropic surfaces, and forms mainly by quasi-horizontal motion. Air with low Q is drawn from low latitudes on the east side of the developing front, cutting across the plane represented by the figure.

The folds shown in Fig.9 and associated atmospheric structures resemble those found near fronts observed in the upper troposphere (e.g. Uccellini et al., 1985, Fig.10), though the patterns are on a much larger scale in the stratosphere. The dynamics of formation of the two types of front may differ, however. For the upper troposphere, a two-dimensional numerical model reproduces some of the features of observed fronts and folds (Hoskins, 1972). But for the stratosphere, results from our model indicate that advection of Q in three dimensions must be represented. Having two examples of a phenomenon in different parts of the atmosphere should further our understanding of both.

4. CONCLUDING REMARKS

Despite the improved coverage that observations from satellites bring, many diagnostic studies of the dynamics of the middle atmosphere begin by zonally averaging such data. One motivation for this is provided by the theory of wave, mean-flow interaction. It contains a body of results on which arguments of cause and effect can be based if departures from some average state are assumed to be small. This assumption is not always valid, however, as synoptic maps for the major warming of 1984/85 vividly illustrate. What constitutes a so-called "pre-conditioned" state for a major warming according to quasi-linear theory is at odds with the actual state of the stratosphere when this warming started to develop rapidly.

Instruments on satellites make measurements of uniform quality, a boon when directly measured quantities must be differentiated one or more times to obtain fields of dynamical interest. But because of the spatial averaging inherent in the measurements, we obtain only a blurred view of the formation of small-scale features--like stratospheric "fronts"-- that might affect the evolution of the large-scale flow. Lidars, radiosondes and rocketsondes measure temperature with much better resolution, but observations are unevenly distributed, if they exist at all. Different types of observations are utilised in weather forecasting by assimilating them regularly into a general circulation model. Dynamically consistent analyses of the state of the troposphere and lower stratosphere are produced as a by-product, including fields of vertical velocity, a quantity not normally measured. It is desirable to extend this procedure to include the whole of the middle atmosphere. Groups involved with the Upper Atmosphere Research Satellite (scheduled for launch in 1991) are planning to do so.

ACKNOWLEDGEMENTS

We thank David Andrews and John Pyle for their helpful comments on this paper. Betty Kingston prepared the diagrams.

REFERENCES

Butchart, N., Clough, S.A., Palmer, T.N. and Trevelyan, P.J. 1982. Simulations of an observed stratospheric warming with quasigeostrophic refractive index as a model diagnostic. Quart. J. R. Met. Soc., 108, 475-502.

Charney, J.G. and Drazin, P.G. 1961. Propagation of planetary-scale disturbances from the lower into the upper atmosphere. J. Geophys. Res., 66, 83-109.

Clough, S.A., Grahame, N.S. and O'Neill, A. 1985. Potential vorticity in the stratosphere derived using data from satellites. Quart. J. R. Met. Soc., 111, 335-358.

Dritschel, D.G. 1986. The nonlinear evolution of rotating configurations of uniform vorticity. J. Fluid. Mech., 172, 157-183.

Holton, J.R. 1976. A semi-spectral numerical model for wave, mean-flow interactions in the stratosphere: application to sudden stratospheric warmings. J. Atmos. Sci., 33, 1639-1649.

Hoskins, B.J. 1972. Non-Boussinesq effects and further development in a model of upper tropospheric frontogenesis. Quart. J. Roy. Met. Soc., 98, 532-541.

Hoskins, B.J., McIntyre, M.I. and Robertson, A.W. 1985. On the use and significance of isentropic potential vorticity maps. Quart. J. R. Met. Soc., 111, 877-946.

Kanzawa, H. 1984. Four observed sudden warmings diagnosed by the Eliassen-Palm flux and refractive index. In Advances in Earth and Planetary Sciences: Dynamics of the Middle Atmosphere, pp. 307-3 . Terra Scientific Publishing Company, Tokyo; D. Reidel Publishing Company, Dordrecht.

Labitzke, K., Naujokat, B., Lenschow, R., Petzoldt, K., Rajewski, B. and Wohlfart, R.C. 1985.. The third winter of MAP-DYNAMICS, 1984/85: a winter with an extremely intense and early major warming. Beilage zur Berliner Wetterkarte, Free University of Berlin.

McIntyre, M.E. and Palmer, T.N. 1983. Breaking planetary waves in the stratosphere. Nature, 305, 593-600.

McIntyre, M.E. and Palmer, T.N. 1984. The 'surf-zone' in the stratosphere. J. Atmos. Terr. Phys., 46, 825-849.

McIntyre, M.E. and Palmer, T.N. 1985. A note on the general concept of wave breaking for Rossby and gravity waves. Pure and Appl. Geophys., 123, 964-975.

O'Neill, A. and Pope, V.D. 1987. Simulations of linear and nonlinear disturbances in the stratosphere. Submitted to Quart. J. Roy. Met. Soc..

Palmer, T.N. 1981. Diagnostic study of a wavenumber-2 stratospheric warming in a transformed Eulerian-mean formalism. J. Atmos. Sci., 38, 844-855.

Pick, D.R. and Brownscombe, J.L. 1981. Early results based on the stratospheric channels of TOVS on the TIROS-N series of operational satellites. Int. Counc. Sci. Unions, Comm. Space Res., Adv. Space Res., 1, No.4, 247-260.

Uccellini, L.W., Keyser, D., Brill, K.F. and Wash, C.H. 1985. The Presidents' Day Cyclone of 18-19 February 1979: Influence of upstream trough amplification and associated tropopause folding on rapid cyclogenesis. Mon. Wea. Rev., 113, 962-988.

FIGURE LEGENDS

Fig.1. (a) Cross-section of zonal-mean, geostrophic wind for the northern hemisphere on 20 December 1984. Contours are plotted at intervals of $10ms^{-1}$, and shading denotes easterly winds. (b) The same but for 2 January 1985.

Fig.2. Synoptic map of Ertel's potential vorticity, Q, and winds on the 850K isentropic surface, near 10mb, for 22 December 1984. Contours of Q are plotted at intervals of 10^{-4} K m^2 kg^{-1} s^{-1}. In these units, dark shading denotes values between 5 and 3; light shading denotes values between 2 and 1. Wind speeds can be read using the scale at the bottom left of the figure.

Fig.3. As for Fig.2 but for 26 December 1984.

Fig.4. As for Fig.2 but for 30 December 1984.

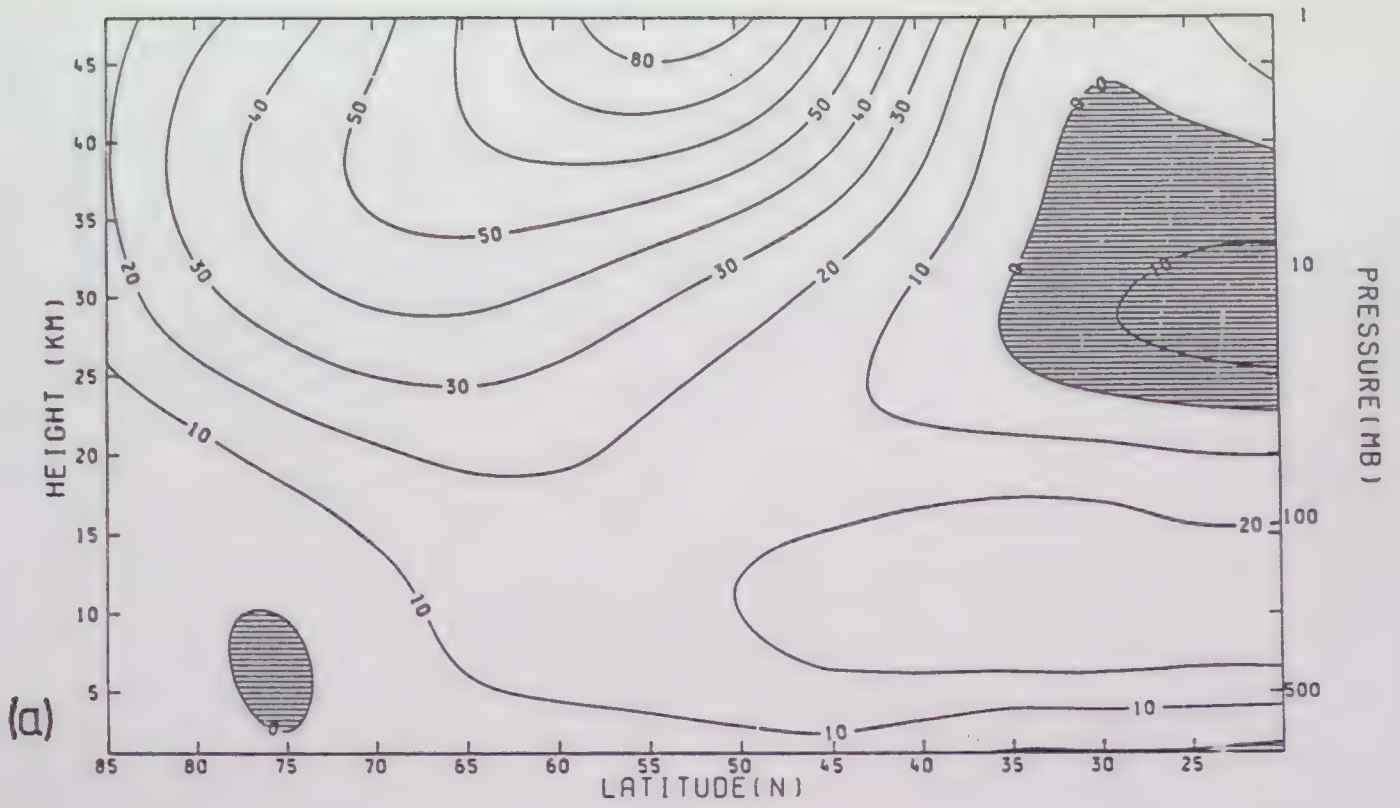
Fig.5. As for Fig.2 but for 2 January 1985.

Fig.6. Co-located profiles of temperature (°C) on 2 January 1985. The solid curve is plotted from measurements made by a radiosonde launched from Berlin. The dashed curve is derived by differentiating in the vertical fields of geopotential height obtained from SSU data.

Fig.7. As for Fig.2 but for simulated fields for 30 December 1984. The numerical experiment started from the observed state of the stratosphere on 22 December 1984.

Fig.8. Longitude-height cross-section at 57.5°N of simulated temperature on 31 December 1984. Contours are plotted at 10K intervals.

Fig.9. Longitude-height cross-section at 57.5°N of simulated ln(Q) on 31 December 1984. Q is converted to units of 10^{-6} K m^2 kg^{-1} s^{-1} before taking logarithm.


(a)
HEIGHT (KM)
LATITUDE (N)
PRESSURE (MB)
80
50
40
30
20
10
0
1
10
100
500
85
80
75
70
65
60
55
50
45
40
35
30
25
5
15
20
25
35
40
45

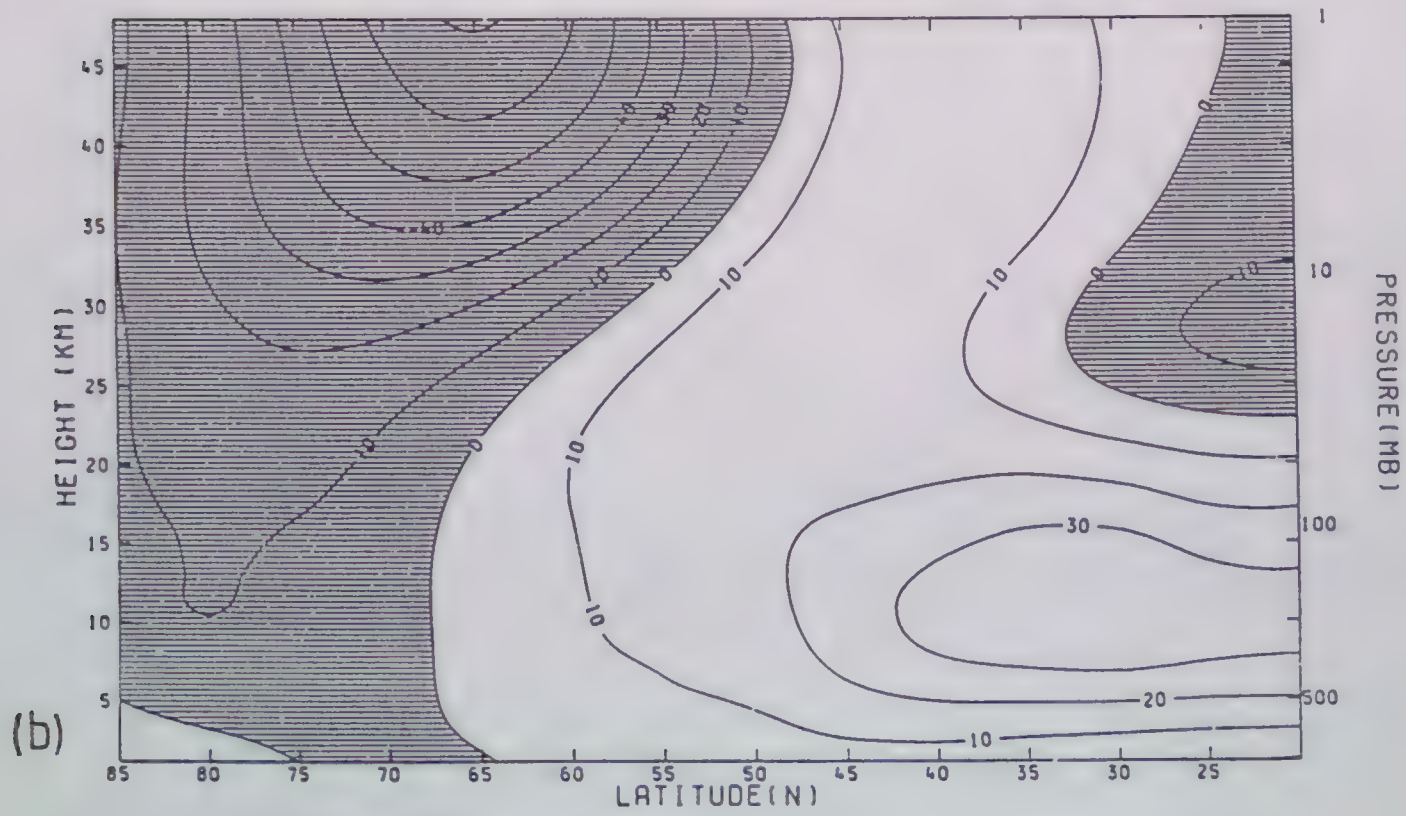

(b)
HEIGHT (KM)
LATITUDE (N)
PRESSURE (MB)
40
30
20
10
0
1
10
100
500
85
80
75
70
65
60
55
50
45
40
35
30
25
5
15
20
25
35
40
45

Fig. 1

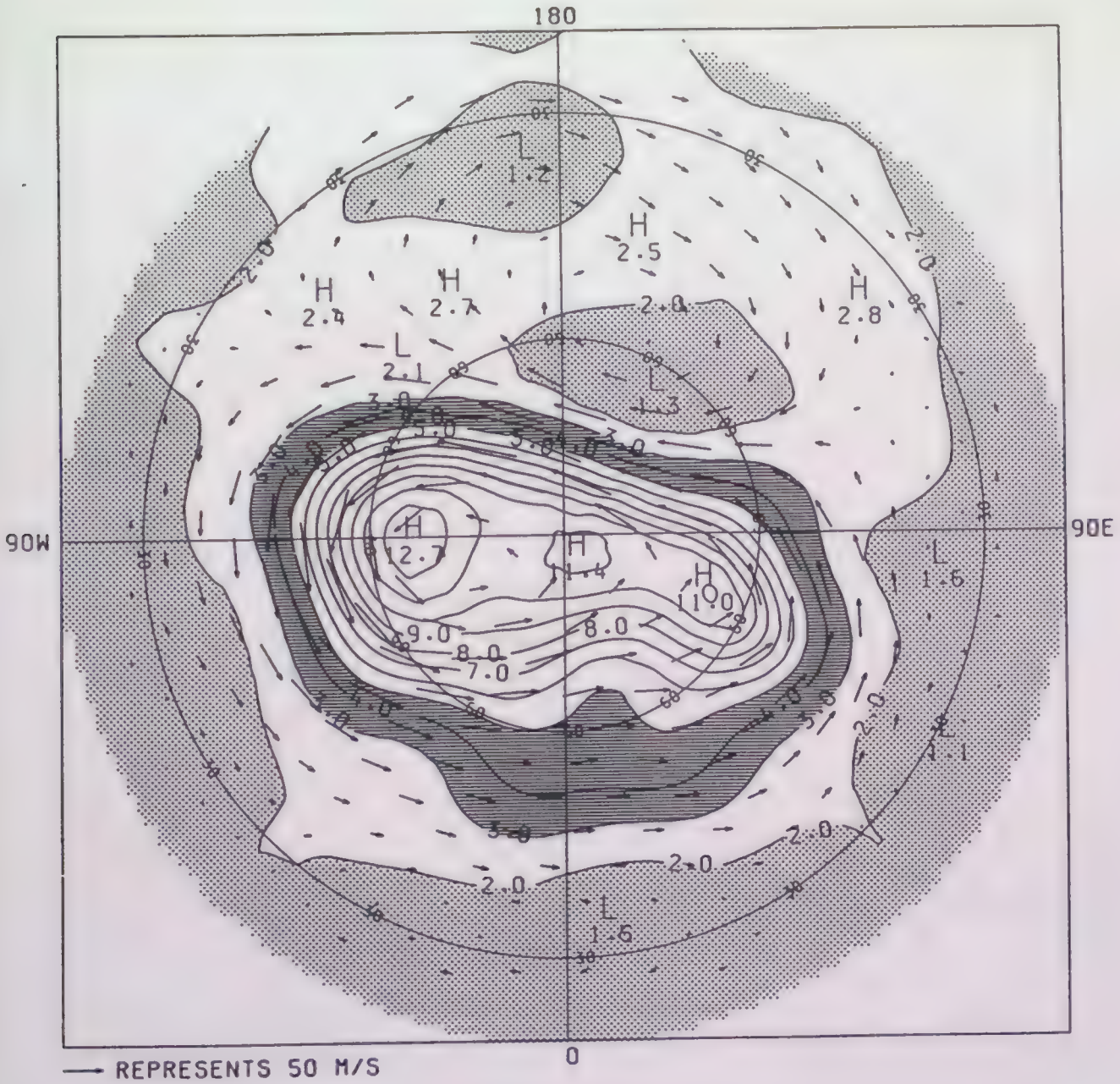

180
90W
90E
0
L
1.2
H
2.5
H
2.4
H
2.7
H
2.8
L
2.1
L
1.3
H
12.7
H
11.4
H
11.0
L
1.6
L
1.1
L
1.6
9.0
8.0
7.0
2.0
3.0
4.0
5.0
REPRESENTS 50 M/S

Fig. 2

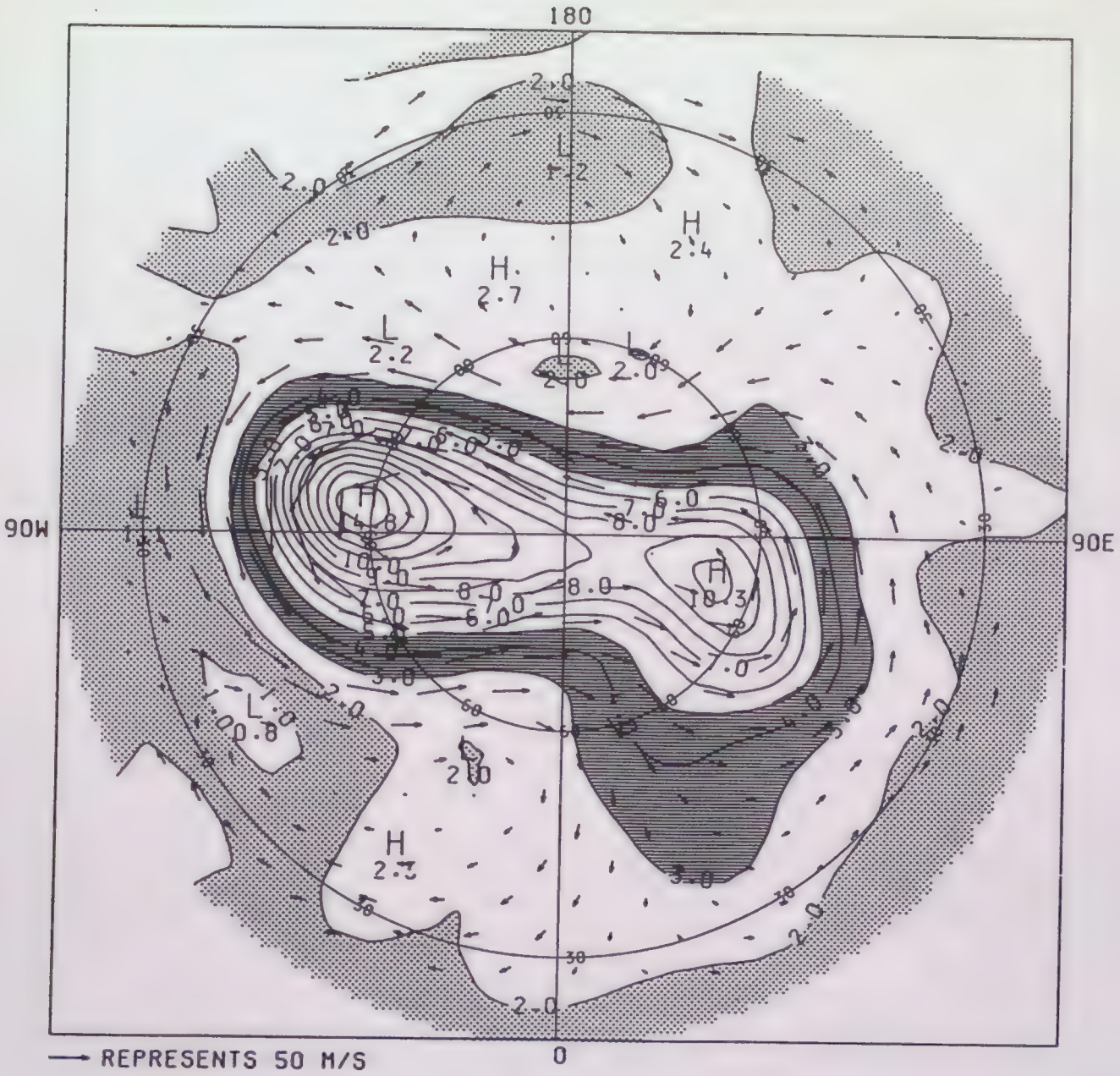


Fig. 3

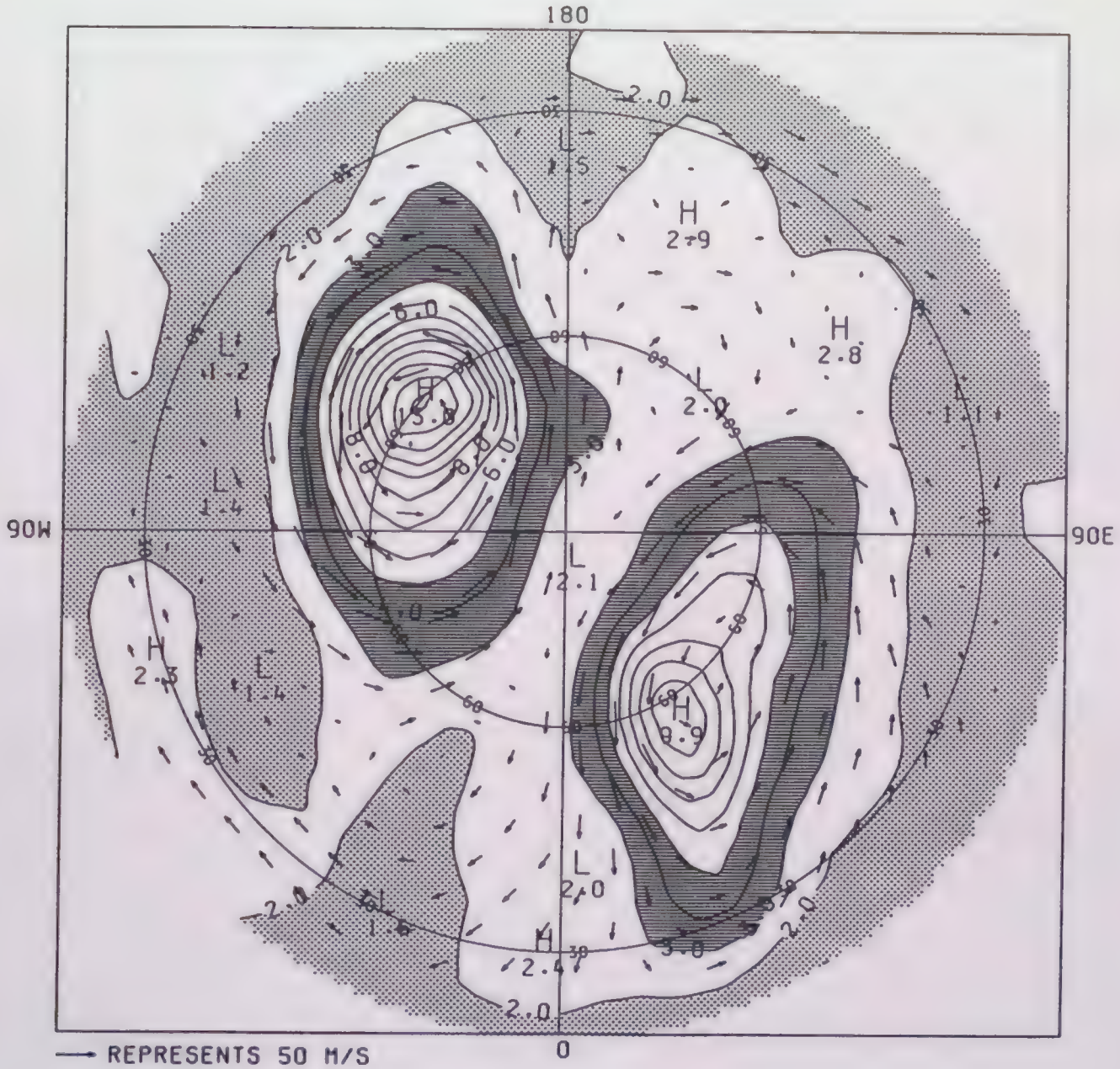


Fig. 4

—— REPRESENTS 50 M/S

Fig. 5

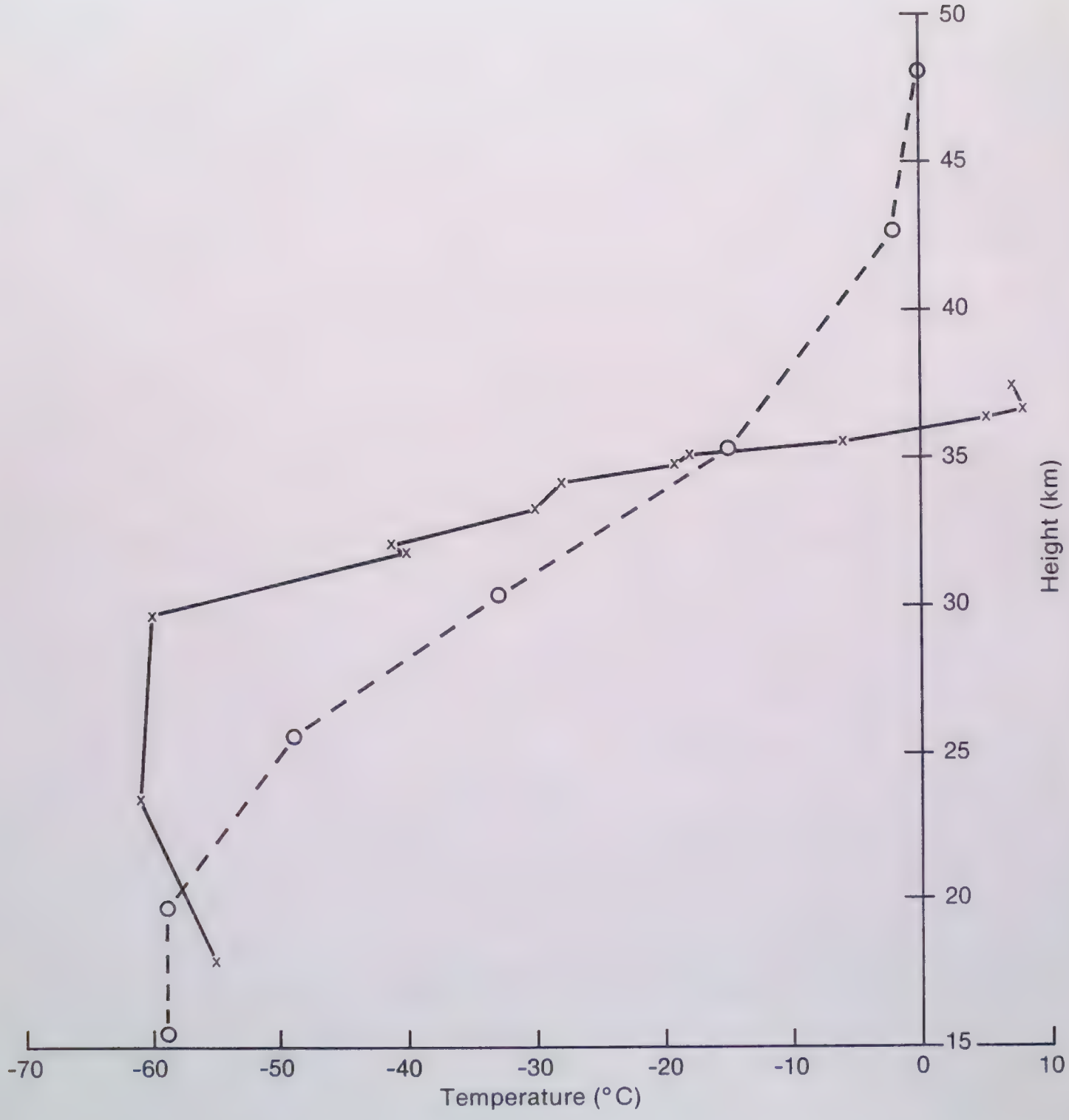

50
45
40
35
30
25
20
15
Height (km)
-70
-60
-50
-40
-30
-20
-10
0
10
Temperature (°C)

Fig. 6

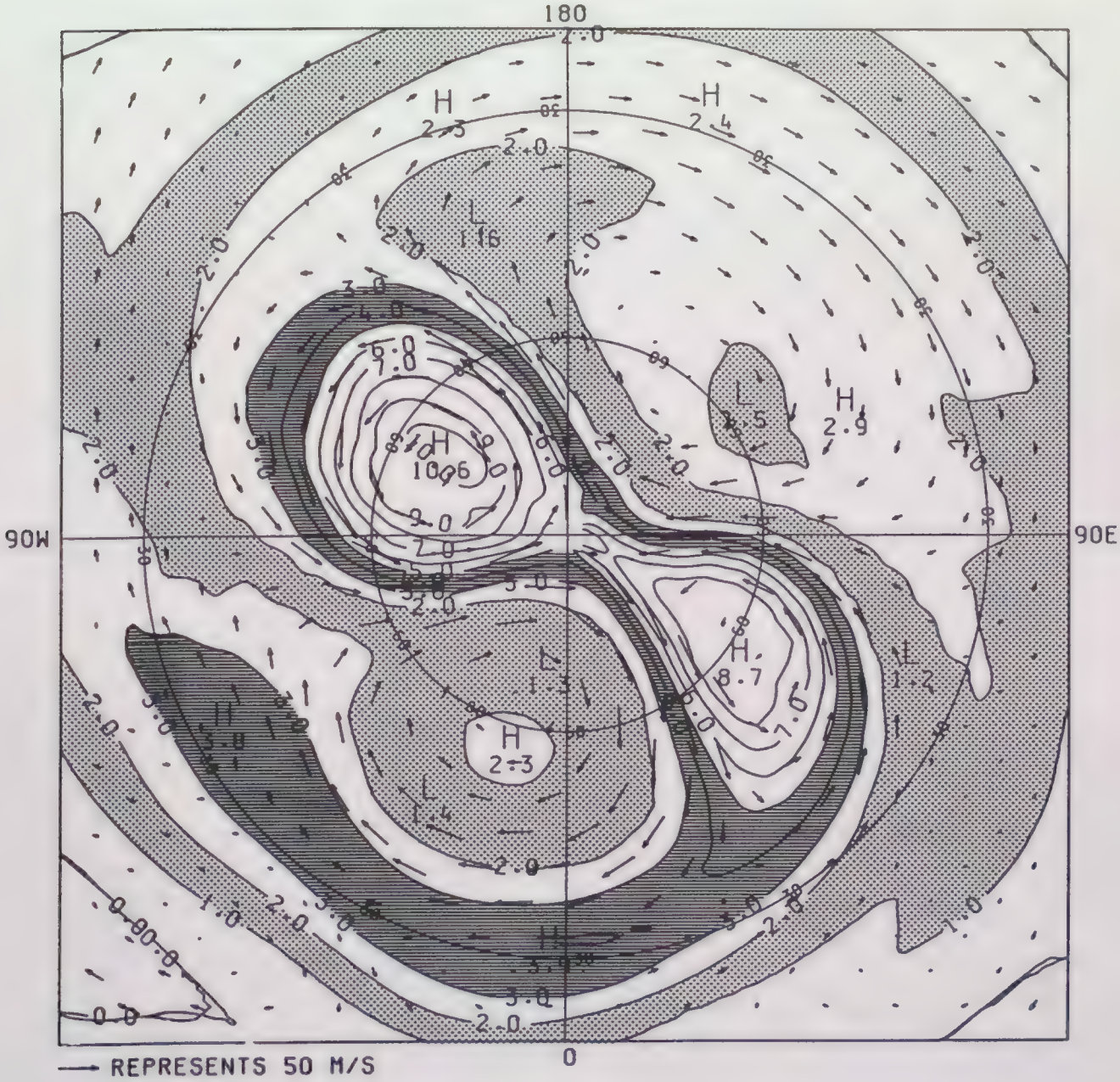


Fig. 7

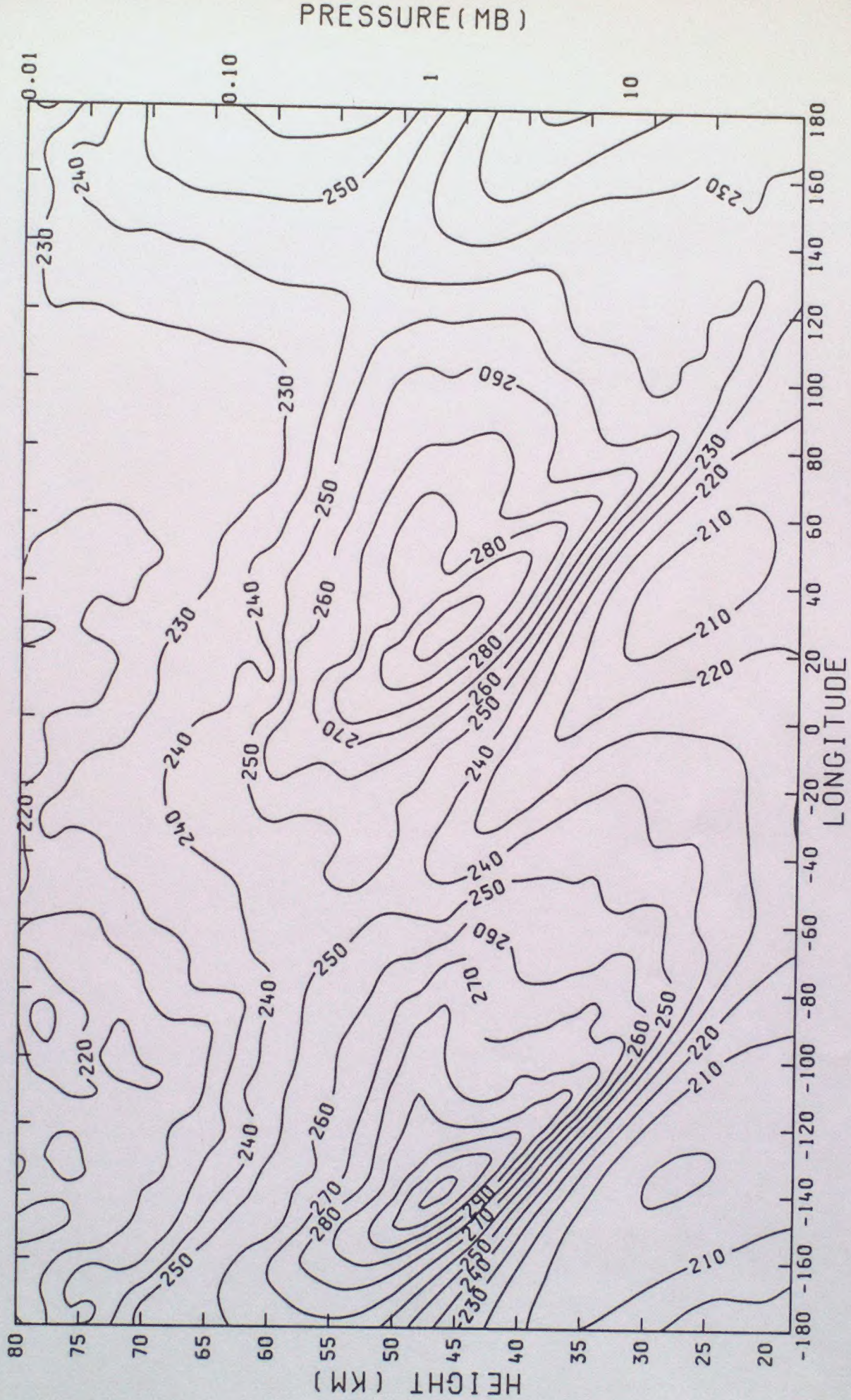

PRESSURE (MB)
0.01
0.10
1
10
HEIGHT (KM)
80
75
70
65
60
55
50
45
40
35
30
25
20
LONGITUDE
-180 -160 -140 -120 -100 -80 -60 -40 -20 0 20 40 60 80 100 120 140 160 180
210
220
230
240
250
260
270
280
290

Fig. 8

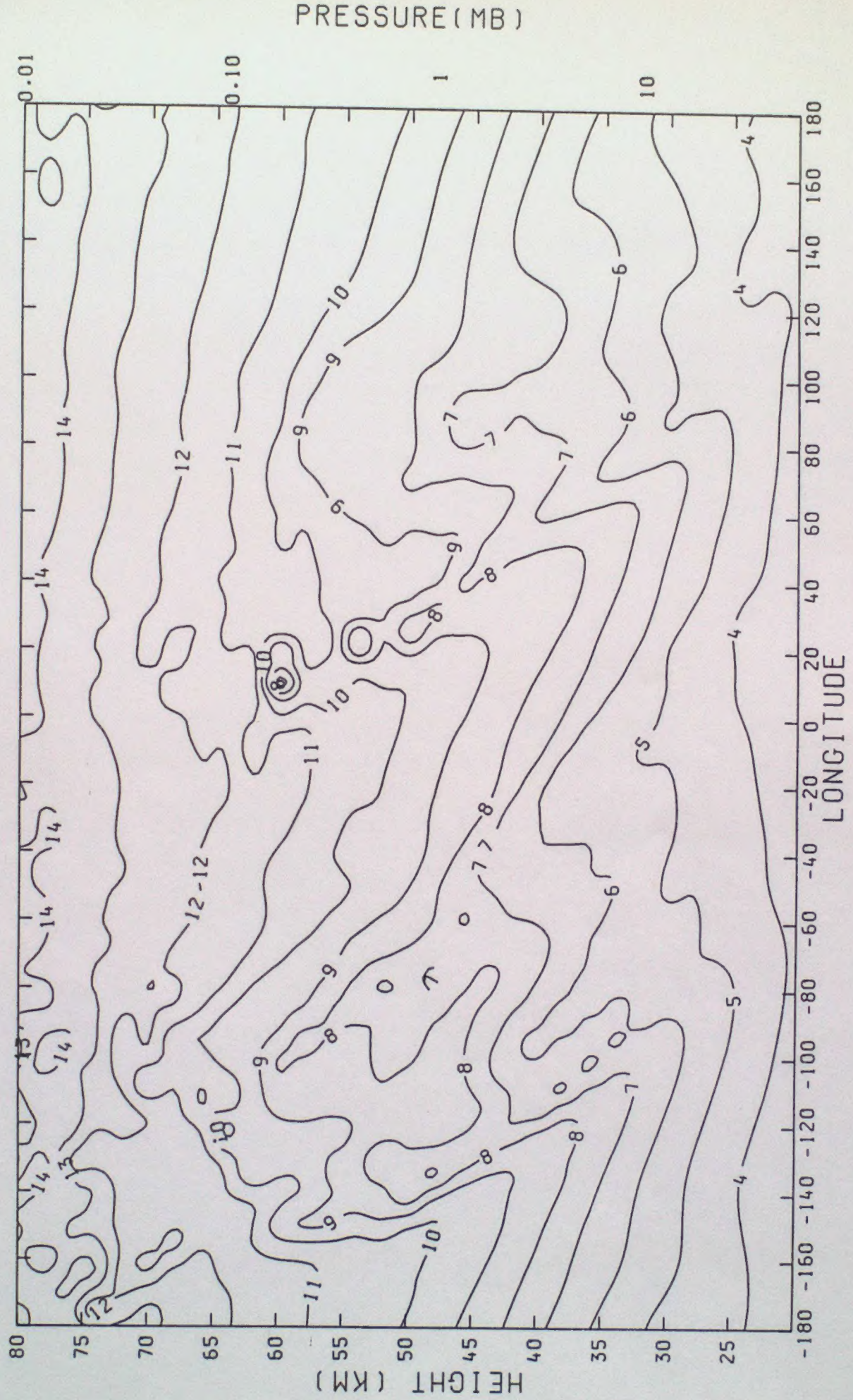

PRESSURE (MB)
0.01
0.10
1
10
HEIGHT (KM)
80
75
70
65
60
55
50
45
40
35
30
25
LONGITUDE
-180 -160 -140 -120 -100 -80 -60 -40 -20 0 20 40 60 80 100 120 140 160 180

Fig. 9

DYNAMICAL CLIMATOLOGY TECHNICAL NOTES

DCTN 1	June 1984	A scheme for incorporating the thermodynamic sea-ice model into a coupled ocean-atmosphere model. C Gordon and M Bottomley
DCTN 2	Sept 1984	The parametrization of the upper ocean mixed layer in coupled ocean/atmosphere models. C Gordon and M Bottomley
DCTN 3	Sept 1984	The response of a simple thermodynamic sea-ice model to GCM forcing appropriate for perennial sea-ice. M Bottomley
DCTN 4	Sept 1984	Potential vorticity in the stratosphere derived using data from satellites. S A Clough, N S Grahame and A O'Neill
DCTN 5	Sept 1984	Simulation of the seasonal cycle of SST using globally specified climatological data to force simple ocean models. N S Grahame
DCTN 6	Sept 1984	Development of an improved GCM radiation scheme. A Slingo and R C Wilderspin
DCTN 7	Sept 1984	On the specification of surface fluxes in coupled atmosphere/ocean general circulation models. J F B Mitchell, C A Wilson and C M Price
DCTN 8	Sept 1984	A survey of possible repercussions of the Soviet river reversal programme on Arctic sea-ice. H Cattle
DCTN 9	Oct 1984	A study of a strong stratospheric warming during January 1982 T D A Fairlie
DCTN 10	Nov 1984	An intercomparison of the Met O 11 M45 model and the Met O 20 11-level model on the Cyber 205 J F Dyson
DCTN 11	Nov 1984	Simulation of observed time series of temperature and precipitation over eastern England. D M Reed
DCTN 12	Nov 1984	Forecasts and analyses of the onset of the Southwest monsoon. R Kershaw
DCTN 13	Nov 1984	On modelling the effects of CO_2 on climate. J F B Mitchell
DCTN 14	Dec 1984	Development of a revised longwave radiation scheme for an atmospheric general circulation model. A Slingo and R C Wilderspin
DCTN 15	Jan 1985	Rossby wavetrains in the stratosphere forced by localised disturbances in the troposphere. C J Marks (Oxford), A O'Neill and V D Pope

DYNAMICAL CLIMATOLOGY TECHNICAL NOTES

DCTN 16	Jan 1985	Observing system experiments using the Meteorological Office's 15-level model. M A Bromley
DCTN 17	Feb 1985	Atmospheric general circulation simulation of the climate of the Mediterranean region. H Cattle
DCTN 18	Feb 1985	Diurnal variation and cloud in a general circulation model. C A Wilson and J F B Mitchell
DCTN 19	Feb 1985	The sensitivity of the Saharan region in a general circulation model. W M Cunnington and P R Rowntree
DCTN 20	Mar 1985	The impact of aircraft observations on analyses and forecasts from the operational 15-level model. B R Barwell
DCTN 21	Mar 1985	Non-linear effects on heat-induced circulation of the tropical atmosphere. A E Gill and P J Phlips
DCTN 22	Apr 1985	Simulation of cloudiness with the U.K. Meteorological Office 11-layer AGCM. A Slingo and R C Wilderspin
DCTN 23	Apr 1985	An integrated approach to the representation of cloud processes. R N B Smith
DCTN 24	May 1985	Nuclear Winter: Calculations of the shortwave radiative effects of soot aerosols. A Slingo and P Goldsmith
DCTN 25	Jul 1985	Sea ice and the Antarctic winter circulation: a numerical experiment. J F B Mitchell and T S Hills
DCTN 26	Jul 1985	A Budyko-Sellers model for study of ocean-atmosphere interaction. W J Ingram
DCTN 27	Aug 1985	The effect of a shoaling thermocline on equatorially-trapped Kelvin waves. A E Gill and B A King
DCTN 28	Sept 1985	CO_2 climate sensitivity and model dependence of results. J F B Mitchell, C A Wilson and W M Cunnington
DCTN 29	Sept 1985	Handbook of the Meteorological Office 11-layer atmospheric general circulation model. Volume 1 : Model Description Edited by A Slingo
DCTN 30	Sept 1985	Impact of land surface variations on African rainfall in general circulation models. P R Rowntree, M F Wilson and A B Sangster

DYNAMICAL CLIMATOLOGY TECHNICAL NOTES

DCTN 31 Nov 1985 The annual cycle of sea level in the eastern tropical Pacific.
G R Bigg and A E Gill

DCTN 32 Nov 1985 An analytic model of the heat-induced tropical circulation in the presence of a mean wind.
P J Phlips and A E Gill

DCTN 33 Nov 1985 The effect of resolution and diffusion on the simulated climate.
J F Dyson

DCTN 34 Jan 1986 Rossby adjustment over a step.
M K Davey, A E Gill, E R Johnson and P F Linden

DCTN 35 Jan 1986 Issues concerning the evaluation of effective surface roughness of heterogeneous terrain.
D J Carson

DCTN 36 Mar 1986 Climate sensitivity and past climates : evidence from numerical studies
J F B Mitchell

DCTN 37 Mar 1986 Parametrizations of land-surface processes in Meteorological Office numerical weather prediction and climate models
D J Carson

DCTN 38 Apr 1986 Modelling of land-surface processes and their influence on European Climate
D A Warrilow, A B Sangster and A Slingo

DCTN 39 Apr 1986 Simulation of climatic change due to increased CO_2
J F B Mitchell

DCTN 40 May 1986 On the interpretation of the Eliassen-Palm flux divergence.
Part I : Theory
D G Andrews

DCTN 41 June 1986 Experiments on tropical circulation with a simple moist model.
M K Davey and A E Gill

DCTN 42 Jun 1986 Parameter regimes for studying isolated eddies.
P D Killworth and M K Davey

DCTN 43 Jun 1986 A comparison of the impact of an envelope orography and of a parametrization of orographic gravity-wave drag on model simulations.
A Slingo and D W Pearson

DCTN 44 Jul 1986 The effect of a sea-surface temperature anomaly on a prediction of the onset of the south-west monsoon over India.
R Kershaw

DCTN 45 Jul 1986 An introduction to the parametrization of land-surface processes.
D J Carson

DYNAMICAL CLIMATOLOGY TECHNICAL NOTES

DCTN 46 Sept 1986 Lateral boundaries and the heat-induced flow in the tropical atmosphere.
P J Phlips and A E Gill

DCTN 47 Nov 1986 Simulation of the diurnal cycle of outgoing longwave radiation with an atmospheric GCM.
A Slingo, R C Wilderspin and S J Brentnall

DCTN 48 Feb 1987 A 2 x CO_2 climate sensitivity experiment with a global climate model including a simple ocean
C A Wilson and J F B Mitchell

DCTN 49 Feb 1987 Sensitivity of the Met O 20 11-layer model to increased land surface roughness length
C C Heasman

DCTN 50 Feb 1987 The influence of atmospheric waves on the general circulation of the middle atmosphere
D G Andrews

DCTN 51 Feb 1987 Transport mechanisms in the middle atmosphere: an introductory survey
D G Andrews

DCTN 52 Mar 1987 The Met O 20 stratosphere-mesosphere model
M Fisher

DCTN 53 Apr 1987 Simulations of linear and nonlinear disturbances in the stratosphere
A O'Neill and V D Pope

DCTN 54 Apr 1987 Aspects of dynamics of the middle atmosphere inferred using data from a satellite and from a numerical model
T D A Fairlie and A O'Neill